1. Laminar flow in a water channel

Balance between gravitational force and vicous stress:

$$u\left(z\right) = \Gamma Hz - \frac{1}{2}\Gamma z^{2}$$
 where $\Gamma = g \cdot \frac{\sin \alpha}{\nu}$

Discharge of water:

$$Q = \int_{0}^{H} \rho \cdot u(z) dz$$

(a) What is the percentage change in depth, if the discharge is increased by 20%: First derive the discharge:

$$Q = \int_0^H \rho \left[\Gamma H z - \frac{1}{2} \Gamma z^2 \right] dz$$
$$= \rho \left[\frac{1}{2} \Gamma H^3 - \frac{1}{6} \Gamma H^3 \right]$$
$$= \frac{1}{3} \rho \Gamma H^3$$

Get now the change in discharche by 20%:

$$Q' = 1.2 \cdot Q$$

$$\frac{1}{3}\rho\Gamma H'^{3} = 1.2 \cdot \frac{1}{3}\rho\Gamma H^{3}$$

$$H'^{3} = 1.2 \cdot H^{3}$$

$$H' = \sqrt[3]{1.2}$$

$$H' \sim 1.063 H$$

The height of the water channel increases by 6.3%.

(b) The change of depth is neither dependent upon the viscosity nor the temperature. Therefore the 6.3% change is valid whatever the fluid properties are.

2. Water flow

One solution of the governing equations is:

$$u\left(z\right) = U\left(1 - e^{-\frac{z}{\varepsilon}}\right)$$

Boundary layer: Layer over which the effects of the surface (e.g. frictional effects) can be felt.

$$\epsilon = \frac{\nu}{W} \qquad \text{where } \begin{cases} \nu \text{ dynamical viscosity} \\ W = 10^{-2} \text{ms}^{-1} \end{cases}$$

	$\nu \ [{ m m}^2 { m s}^{-1}]$	$\varepsilon \ [\mathrm{m}]$
Water	$1.1 \cdot 10^{-6}$	$1.1 \cdot 10^{-4} = 0.1 \text{ mm}$
Air	$1.5 \cdot 10^{-5}$	$1.5 \cdot 10^{-3} = 1.5 \text{ mm}$

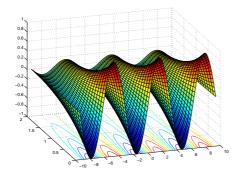
So the boundary layer of the atmosphere is generally speaking much deeper than that of the ocean.

3. 2D flow field

Streamfunction:

$$\psi = A\sin(kx) e^{-ly}$$

(a) Streamfunction and streamlines



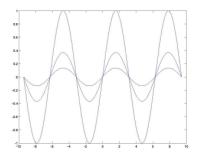
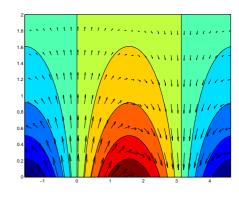


Figure 1: Streamfunction (left) and the some corresponding streamlines (right).

(b) Velocity components u and v:

$$u = -\frac{\partial \psi}{\partial y} = Al \sin(kx) e^{-ly}$$
$$v = -\frac{\partial \psi}{\partial x} = Ak \cos(kx) e^{-ly}$$



(c) Vertical vorticity:

$$\zeta = \frac{\partial v}{\partial x} - \frac{\partial u}{\partial y}$$

$$\frac{\partial v}{\partial x} = -Ak^2 \sin(kx) e^{-ly}$$

$$\frac{\partial u}{\partial y} = -Al^2 \sin(kx) e^{-ly}$$

$$\zeta = (l^2 - k^2) A \sin(kx) e^{-ly} = (l^2 - k^2) \psi$$

(d) Values of k and l to be a solution of:

$$\frac{D\zeta}{Dt} = \nu \cdot \nabla^2 \zeta$$

$$\frac{D}{Dt}(l^2 - k^2)\psi = \nu \nabla^2 (l^2 - k^2)\psi$$
$$(l^2 - k^2) \left[\frac{D\psi}{Dt} - \nu \nabla^2 \psi \right] = 0$$

Two possibilities:

$$(l^{2} - k^{2}) = 0$$

$$l = \pm k$$
or
$$\frac{D\psi}{Dt} - \nu \nabla^{2} \psi = 0$$

$$k = 0 \text{ (trivial solution)}$$

4. Shallow water system

Set of equations (6.16) in lecture notes on page 65. 1-dimensional variation in y, z direction $\frac{\partial}{\partial y}, \frac{\partial}{\partial z} = 0$.

(a) Normalisation and parameter \exists

$$\frac{\partial u}{\partial t} + u \frac{\partial u}{\partial x} - fv = -g \frac{\partial h}{\partial x}$$
 (1)

$$\frac{\partial v}{\partial t} + u \frac{\partial v}{\partial x} - f u = 0 \tag{2}$$

$$\frac{\partial h}{\partial t} + u \frac{\partial h}{\partial x} + h \frac{\partial u}{\partial x} = 0 \tag{3}$$

Normalize with: $\tilde{u}=\frac{u}{U},\,\tilde{h}=\frac{h}{H},\,\tilde{x}=\frac{x}{L},\,\tilde{t}=\frac{U}{L}t$

• (1)

$$\frac{\partial u}{\partial t} = \frac{U^2}{L} \frac{\partial \tilde{u}}{\partial \tilde{t}}, \qquad u_t = \frac{U^2}{L} \tilde{u}_{\tilde{t}}$$

$$\frac{\partial u}{\partial x} = \frac{U}{L} \frac{\partial \tilde{u}}{\partial \tilde{x}}, \qquad u_x = \frac{U}{L} \tilde{u}_{\tilde{x}}$$

$$\frac{\partial h}{\partial x} = \frac{H}{L} \frac{\partial \tilde{h}}{\partial \tilde{x}}, \qquad h_x = \frac{H}{L} \tilde{h}_{\tilde{x}}$$

$$\frac{U^2}{L}\left(\tilde{u}_{\tilde{t}} + \tilde{u}_{\tilde{x}}\right) - fU\tilde{v} = -g\frac{H}{L}\tilde{h}_{\tilde{x}}$$

With $R_0 = \frac{U}{fL}$ follows:

$$R_0 \left(U f \left(\tilde{u}_{\tilde{t}} + \tilde{u}_{\tilde{x}} \right) \right) - f U \tilde{v} = R_0 \left(-g f \frac{H}{U} \right) \tilde{h}_{\tilde{x}}$$

$$R_0 \left(\tilde{u}_{\tilde{t}} + \tilde{u}_{\tilde{x}} \right) - \tilde{v} = R_0 \left(-g \frac{H}{U^2} \right) \tilde{h}_{\tilde{x}}$$

Therefore: $\exists = -g\left(\frac{H}{U^2}\right)$

(2)

$$\frac{U^2}{L}\tilde{v}_{\tilde{t}} + \frac{U^2}{L}\tilde{u}\tilde{v}_{\tilde{x}} + fU\tilde{u} = 0$$

$$R_0 (fU\tilde{v}_{\tilde{t}} + fU\tilde{u}\tilde{v}_{\tilde{x}}) + fU\tilde{u} = 0$$

$$R_0 (\tilde{v}_{\tilde{t}} + \tilde{u}\tilde{v}_{\tilde{x}}) + \tilde{u} = 0$$

• (3)

$$\frac{UH}{L}\tilde{h}_{\tilde{t}} + \frac{UH}{L}\tilde{h}_{\tilde{x}}\tilde{u} + \frac{UH}{L}\tilde{u}_{\tilde{x}}\tilde{h}u = 0$$
$$\tilde{h}_{\tilde{t}} + \tilde{h}_{\tilde{x}}\tilde{u} + \tilde{u}_{\tilde{x}}\tilde{h}u = 0$$

(b) Characteristic values

$$\begin{array}{rcl} R_0 & = & \frac{U}{fL} = \frac{10^{-2}\,\mathrm{ms}^{-1}}{10^{-4}\,\mathrm{s}^{-1}10^4\,\mathrm{m}} = 10^{-2} \\ & \to & \text{Corioliseffect is important} \\ \exists & = & -\frac{gH}{U^2} = \frac{10\,\mathrm{ms}^{-2}\cdot10\,\mathrm{m}}{\left(10^{-2}\,\mathrm{ms}^{-1}\right)^2} = 10^6 \end{array}$$

(c) Dynamical similarity

The combination of some key elements is more important than elements themselves separately. Different configuration but same combination \Rightarrow same flow response similar for same \exists , R_0 as (b).

Assumptions: $R_0 = 10^{-2}$, $\exists = 10^6$ Lake: $U_L = 10^{-2} \text{ ms}^{-1}$, $L_L = 10^4 \text{ m}$ Model: $U_M = 10^{-3} \text{ ms}^{-1}$, $L_M = 1 \text{ m}$

Depth $H_{\text{Model}}, f_{\text{Model}} = ?$

$$f_{\text{Model}} = \frac{U_M}{R_0 L_M} = \frac{10^{-3} \,\text{ms}^{-1}}{10^{-2} \cdot 1 \,\text{m}} = 10^{-1} \,\text{s}^{-1}$$

Since
$$f = 2\omega \Rightarrow \omega = 5 \cdot 10^{-2} \text{s}^{-1}$$

 $\Rightarrow T = \frac{2\pi}{\omega} \approx 126 \text{s} \sim 2 \text{min}$

$$H_{\text{Model}} = \frac{\exists U_M^2}{g} = \frac{10^6 \cdot (10^{-3} \text{ms}^{-1})^2}{10 \text{ms}^{-2}} = 10^{-1} \text{m}$$